Research Article

Spatial variability and mapping of selected soil properties in Kaligaudanahalli Microwatershed, Gundlupet Taluk, Chamarajanagar District, under hot semi arid agrosubregion of Central Karnataka Plateau, India



Rajendra Hegde¹ · G. Bardhan¹ · K. V. Niranjana¹ · B. P. Bhaskar¹ · S. K. Singh²

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Abstract

The soil attributes of spatially structured data were often measured in Central plateau, of Karnataka, but often poorly defined spatial variability of soil attributes important for land management under hot semiarid agroecological environment. The Kaligaudanahalli microwatershed (612 ha) divided into 250 m × 250 m square grids using ArcGIS and collected surface soils (0–15 cm) at 89 locations. The results of cluster analysis with average linkage method revealed four soil typologies viz., slightly alkaline, sandy clay loam texture showing high coefficient of variation for all soil variables in cluster-I, but neutral, sandy clay loam with high variability for DTPA-extractable Cu in cluster-III and IV. The study concluded that the measured soil properties in regular gird sampling at given scale were enough to capture spatial dependence using ordinary and cokriging techniques and in deriving thematic maps for efficient soil management strategies at watershed level. The results of spatial dependence of each soil property, using ordinary kriging showed that there is strong spatial dependence of pH, EC, K₂O, B, Cu, Fe and Mn with nugget-to-sill ratio [($C_0/C_0 + C$)] less than 0.25 showing strong spatial distribution where DTPA-extractable Zn had a nugget/sill more than 0.75, displaying a weak spatial autocorrelation. The results showed that cokriging is the best fit and strongly correlated with EC, P₂O₅, K₂O, Fe and Mn using pH as ancillary variable over ordinary kriging to derive thematic maps.

Keywords Cluster analysis · Cokriging · Kriging · Microwatershed · Spatial variability · Semivariogram

1 Introduction

The spatial variability of soil properties can be measured and quantified at a given scale with regular sampling technique for designing sustainable land management practices [19, 43]. Most of the studies on spatial variability of soil attributes using geospatial techniques have been carried out in diverse temperate countries [5, 7, 10, 12, 15, 47]. In the past three decades, the application of geostatistical methods by soil researchers primarily focused on predicting spatial variability along a transect [40] or at field-scale [8, 43]. Using geostatistics, GIS and remote sensing techniques for large area in Australia, McBratney et al. [26] provided the comprehensive maps for physical, chemical and biological soil properties. In India, majority of soil maps were prepared by conventional methods on smaller scale. Due to initiation of detailed microwatershed surveys in drought prone areas of Karnataka along with grid survey, the modern spatial prediction techniques were employed in spatial variability of soil properties [25, 28, 37]. It was reported that soil pH, electrical conductivity, organic carbon and exchangeable bases have been

Rajendra Hegde, rajendrahanagar@gmail.com | ¹National Bureau of Soil Survey and Land Use Planning (ICAR) (NBSSLUP), Hebbal, Bengaluru, India. ²NBSSLUP (ICAR), Amravathi Road, Nagpur 33, India.



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varied highly in acid soils of India [2], the spatial variability of organic carbon, available N, P and K of Karlawad village in Dharwad district of Karnataka [30] and in lateritic soils of West Bengal [4].

Several workers reported that cokriging employs spatial information on the primary variables along with the spatial correlation between primary and auxiliary variables [16, 45, 48, 54]. A recent study [44] made a comparative study of interpolation techniques and confirmed that cokriging is performed well over ordinary kriging. Considering the review of the literature, we hypothesized to answer, what statistical tools are useful in explaining the variability of soil properties and their use in designing crop specific nutrient management plans? The research herein is a part of Sujala-III project in semiarid tracts of Karnataka State for reducing fertilizer recommendation and enhancing crop productivity of small-scale farmers. Hence, the objectives of this study were to (i) evaluate the efficacy of grid survey method in addressing top soil properties and (ii) analyze the variability of physicochemical properties of top soil by means of cluster analysis, and their spatial dependence by ordinary kriging and cokriging comparison through cross-validation.

2 Materials and methods

2.1 Study area

The Kaligaudanahalli microwatershed lies between 11°41'-11°43'N and 76°37'-76°41'E covering 612 ha. It is bounded by Karle, Siddapura, Siddainapura and Maguvinahalli villages in Gundlupet taluk, Chamarajanagar district, Karnataka State, India (Fig. 1). The study area is a part of hot moist semiarid subregion of Central Karnataka Plateau, with medium to deep red loamy soils, low AWC (available water holding capacity) and LGP (length of growing period) of 120–150 days [46]. The elevation is 450–900 m above mean sea level with mean annual rainfall of 670-890 mm of which about 55-75% is received during the south west monsoon. Geological formations in the study area show the occurrence of Archean-age granite gneiss and schistose rocks with quartzite and banded iron formations and Dolerite dykes of Proterozoic age with variable strike directions and width [3]. It was reported that Typic/Lithic Rhodustalfs/Haplustalfs are dominant but in association with Typic/Vertic Haplustepts under isohyperthermic soil temperature regime and ustic soil moisture regime [32].

2.2 Grid soil survey

The grid survey at an interval of by 250×250 m was carried out to collect soil samples of plow layer (0–15 cm). A total of 89 grid points in Kaligaudanahalli microwatershed were taken. At each grid point, 8–10 cores samples within 10 m radius were taken and georeferenced using GPS system (GARMIN etrex-30x model) (Fig. 2).

2.3 Laboratory methods

The collected surface soil samples were air-dried, grinded and passed through 2-mm sieve for fine earth fraction and used fine earth fraction (< 2 mm) for laboratory analysis. Soil reaction was determined in 1:2.5 soil/water suspension using standard pH meter (Hanna instrument) and electrical conductivity (EC) using conductivity meter (Toshniwal insts. Mfg. Pvt. Ltd, TC 17). Soil organic carbon (OC) was estimated using the wet digestion method as described by Walkley and Black (1934) [50] and expressed its value in percentage. The available potassium (K⁺) in 1N ammonium acetate (NH₄OAc) at (pH 7.0) was determined by Flame photometer [17]. DTPA-extractable micronutrients viz., Fe, Mn, Zn and Cu were extracted by 0.005 M DTPA at pH 7.3 according to the method of [23], and the concentration of micronutrients was estimated using atomic absorption spectrophotometer (Agilent Technologies, 200 series AA model). Available phosphorus was determined by Olsen method [27] and available S by CaCl₂ extraction method [53]. The available Boron was estimated by azomethine-H method by using Lab India (analytical) UV/VIS spectrophotometer [14]. Fertility status of N, P, K and S were categorized as low, medium and high and that of DTPA-extractable zinc, iron, copper and manganese interpreted as deficient, sufficient and excess as per the criteria of Arora [1].

2.4 Statistical analysis

2.4.1 Hierarchical cluster analysis and descriptive statistics

Hierarchical cluster analysis was used to categorize the 11 soil properties at 89 locations with similar characteristics into a group using average. CA was performed to identify analogous behavior among the different soil properties and also among the samples. It was performed on the normalized data set by means of Ward's method using squared Euclidean distances as a measure of similarity (SPSS 15 version). The clusterwise descriptive analysis was made and classified on the basis of



Fig. 1 Location map of Kaligaudanahalli microwatershed

coefficient of variation as low: < 15%, moderate: from 15 to 50% and high as > 50% [51]. Correlations and KS test was performed by XLSTAT 20191.1.

2.4.2 Geostatistical analysis

The soil properties were analyzed using geostatistical tool and calculated semivariogram [16, 19] for each soil property using the equation as given below:

$$\gamma(h) = \frac{1}{2N(h)} + \sum_{i=i}^{N(h)} \left[z(x_i + h) - z(x_i) \right]^2,$$
(1)

where γ (*h*) is the experimental semivariogram value at distance interval *h*; *N* (*h*) is the number of sample value pairs within the distance interval *h*; *z* (*xi*), *z* (*xi* + *h*) is sample values at two points separated by the distance interval *h*. Exponential,

$$\gamma(h;\theta) = \left\{ \left\{ c_0^0 + c_1 \left(1 - \exp\left(-\frac{h}{\alpha}\right) \right) \mid h = 0; h \neq 0, \\ \theta = \left(c_{0,}c_1, \alpha \right)', \text{ where } c_0 \ge 0, c_1 \ge 0, \text{ and } \alpha \ge 0. \right.$$
(2)

The parameters of the model: nugget semivariance, range and sill or total semivariance were determined. To

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Fig. 2 Map showing grid points in the Kaligaudanahalli microwatershed

define different classes of spatial dependence for the soil variables, the ratio between the nugget and the sill was used [8] to classify as: $\leq 25\%$ -strongly spatially dependent, between 26 and 75%-moderately spatially dependent, > 75%-weakly spatially dependent and 100% considered as non-spatially correlated (pure nugget). We employed ordinary [13] and cokriging techniques, and used the procedures as given under:

2.4.2.1 Ordinary kriging (OK) Ordinary kriging is an univariate linear regression model which was used to estimate variance and confidence interval assuming a normal distribution of errors is the spatial prediction of the unmeasured point s_0 is given by predicting the value $\hat{z}(s_0)$, which equals the line sum of the known measured values (i.e., observed values). The formula was used to workout spatial variability as given under:

$$\hat{z}(s_0) = \sum_{i=i}^n \lambda_i z(s_i)$$
(3)

where $\hat{z}(s_0)$ is the value at location s_0 to be interpolated, $z(s_i)$ are the sampled values, and λi , determined by the semivariogram modeling and are the weights to be assigned to each unsampled location.

2.4.2.2 Cokriging (CoK) Cokriging was worked out based on spatial structure of co-variables and their covariance (autocorrelation). The cokriging (CoK) estimator ds at the unsampled location x_0 , Z_w cross-correlated with the main variable Z_v , was given as [16]:

$$Z_{v}^{*}(x_{0}) = \sum_{i=1}^{n} \alpha_{i} \cdot Z_{v(x_{i})} + \sum_{j=1}^{m} \beta_{j} \times Z_{w(x_{j})}$$
(4)

where α_i , β_j —weights assigned to the known values of the primary and secondary variables Z_v and $Z_{w'}$ respectively, n, m—numbers of primary and secondary observations.

2.4.3 Cross-validation technique

Cross-validation technique was used to examine the difference between the known data and the predicted data using the mean error [ME, Eq. (5)], the root-mean-squared error [RMSE, Eq. (6)], the average kriging standard error [AKSE, Eq. (7)], the root-mean-square standardized prediction error [RMSP, Eq. (8)] and the mean standardized prediction error [MSPE, Eq. (9)]. Equations (7) and (9) are only applicable to kriging as they require the kriging variance. Equations (5) and (6) are applicable to all of the interpolation techniques applied in this research.

$$\mathsf{ME} = \frac{1}{N} \sum_{i=N}^{N} \{ Z(x_i) - \hat{Z}(x_i) \}$$
(5)

RMSE =
$$\sqrt{\frac{1}{N}} \sum_{i=1}^{N} \{z(x_i) - \hat{z}(x_i)\}^2$$
 (6)

$$\mathsf{AKSE} = \sqrt{\frac{1}{N}} \sum_{i=1}^{N} \sigma^2(x_i) \tag{7}$$

$$\mathsf{RMSP} = \sqrt{\frac{\sum_{i=1}^{n} \left[\left(\hat{Z}s_i \right) - z(s_i) / \hat{\sigma}(s_i) \right]^2}{n}} \tag{8}$$

$$MSPE = \frac{\sum_{i=1}^{n} (\hat{Z}s_i) - z(s_i) / \hat{\sigma}(s_i)}{n}$$
(9)

where $\hat{Z}(s_i)$ is the predicted value, $z(s_i)$ is the observed (known) value, N is the number of values in the dataset, and σ^2 is the kriging variance for location s_i [18, 21, 49, 52].

3 Results

3.1 Cluster analysis and sample grouping

The cluster analysis shows that cluster-II (45% of total samples) has neutral reaction followed by slightly alkaline under cluster-I and III (55.2% of total sample) and moderately alkaline with 9.1% of grids under cluster-IV. The mean pH is 7.78 ± 0.61 with CV of 7.81% for cluster-I and 7.9 ± 0.59 with CV of 7.5% for cluster-III. The cluster-II has mean pH of 6.75 ± 0.90 with CV of 13.4% but for cluster-IV having mean of 8.21 ± 0.26 with CV of 3.2% (Table 1).

These soils are nonsaline with electrical conductivity less than 0.37 dS m⁻¹ and high coefficient of variation in cluster-I and II (CV > 35%). Considering the trends of soil test values [29], these soils have high soil organic carbon with mean of 0.81 ± 0.39 for cluster-I, 0.77 ± 0.25 for cluster-II and 0.83 ± 0.51 for cluster-IV, but medium in case of cluster-III with mean organic carbon content of 0.63 ± 0.28 . The high variability of soil organic carbon is recorded in cluster-I (CV of 48.4%), cluster-II (CV of 43.99%) and cluster-IV (CV of 61.15%), but medium in case of cluster-III with CV of 32.6%. The available P_2O_5 in top soils is categorized as high in cluster-IV (mean of 90.23 ± 43.02 kg ha⁻¹), whereas medium in cluster-I and III and low in cluster-II (mean of 14.06 ± 11.72 kg ha⁻¹). Similarly, the status of available K₂O is high in all clusters with high variability in cluster-II (CV of 35.35%). The sulfur status is medium for cluster-III and IV but low in cluster-I and II. The coefficient of variation is high for all clusters regarding available P_2O_5 and S status (Table 1). The available boron status is more than 0.5 mg kg⁻¹ in all clusters with high CV except in cluster-II where B is medium $(0.44 \pm 0.3 \text{ mg kg}^{-1})$. Among DTPAextractable micronutrients, Cu, Mn and Fe contents shows sufficient in all clusters with high coefficient of variation. It is interesting to note that DTPA-extractable Zn is below critical limit in cluster-II (mean of 0.51 ± 0.37 mg kg⁻¹) and cluster-III (0.56 ± 0.28 mg kg⁻¹). These clusters have CV of 71.86% (cluster-II) and 49.18% (cluster-III).

The Kolmogorov–Smirnov statistic quantifies a distance between the empirical distribution function of the sample and the cumulative distribution function of the reference distribution, or between the empirical distribution functions of two samples and calculates a P value from that and the sample sizes. If the P value is small, conclude that the two groups were sampled from populations with different distributions. The populations may differ in median, variability or the shape of the distribution.

(KS) normality test indicates that all soil variables are very significant (< 0.001) to extremely significant (< 0.0001) in all clusters except two variables. As a result of using software to test for normality, small *p* values in the output generally indicate the data is not from a normal distribution [36]. Kolmogorov–Smirnov (KS) normality test indicates that P₂O₅ is almost normally distributed with a non-significant (high) *p* value of 0.388 in cluster-I, 0.076 in cluster-III, 0.022 in cluster-IV and Mn with high *p* value of 0.007 in cluster-II and cluster-IV. Except above two variables, all soil variables are very significant (< 0.001) to extremely significant (< 0.0001) in all clusters (Table 1).

3.2 Correlation matrix

The correlation analysis shows that pH has significant (1% level) positive correlation with EC (0.707**), P_2O_5 (0.382**),

Table 1 Clusterw	ise descriptive	estatistics of sc	oil properties i	in Kaligaudanaha	Ili microwatershed						
	Hq	EC (dS m^{-1})	OC (%)	P_2O_5 (kg ha ⁻¹)	K_2O (kg ha ⁻¹)	S (mg kg ⁻¹)	B (mg kg ⁻¹)	Cu (mg kg ⁻¹)	Fe (mg kg^{-1})	Zn (mg kg ⁻¹)	Mn (mg kg^{-1})
Cluster-I											
Mean±SD	7.78±0.61	0.21 ± 0.08	0.81 ± 0.39	45.22 ± 50.01	733.44 ± 83.87	19.21±7.41	0.51 ± 0.36	1.97 ± 0.74	7.74 ± 5.11	0.67 ± 0.39	14.16 ± 8.12
C	7.81	35.69	48.40	110.57	11.44	38.56	70.89	37.65	65.93	58.60	57.35
Max	8.65	0.37	1.86	158.67	875.8	31.93	1.67	3.34	19.24	1.75	28.80
Min	6.07	0.11	0.17	1.00	593.8	3.33	0.08	0.82	2.35	0.24	5.25
CL (95.0%)	0.23	0.03	0.15	19.09	32.36	2.80	0.14	0.27	1.84	0.15	3.04
Kurtosis	1.33	- 0.67	0.32	0.35	- 1.12	- 0.70	3.02	- 0.70	0.11	1.33	- 1.25
Skewness	- 1.12	0.42	0.66	1.24	0.00	0.12	1.58	0.44	1.04	1.35	0.54
<i>p</i> value (2-tailed)	< 0.0001	< 0.0001	< 0.0001	0.388	< 0.0001	< 0.0001	< 0.0001	< 0.0001	< 0.0001	< 0.0001	< 0.0001
Cluster-II											
Mean±SD	6.75 ± 0.90	0.14 ± 0.06	0.63 ± 0.28	13.90 ± 11.72	348.0 ± 123.0	16.85 ± 9.15	0.44 ± 0.30	1.59 ± 0.51	9.61 ± 5.05	0.51 ± 0.37	23.95 ± 12.66
C	13.40	41.57	43.99	84.30	35.35	54.28	67.52	31.96	52.56	71.86	52.87
Max	8.60	0.25	1.52	53.10	563.9	35.83	1.07	3.02	20.00	1.68	62.70
Min	5.00	0.06	0.17	1.00	102.4	0.83	0.05	0.68	2.33	0.16	56.85
CL (95.0%)	0.30	0.02	0.09	3.81	39.13	3.00	0.10	0.16	1.66	0.12	4.15
Kurtosis	-0.70	0.80	1.38	1.88	- 0.90	- 0.78	- 0.66	0.35	- 0.51	1.84	0.67
Skewness	0.35	0.90	1.05	1.34	0.01	0.05	0.57	0.35	0.57	1.53	0.58
<i>p</i> value (2-tailed)	< 0.0001	< 0.0001	< 0.0001	< 0.0001	< 0.0001	< 0.0001	< 0.0001	< 0.0001	< 0.0001	< 0.0001	0.007
Cluster-III											
Mean±SD	7.90 ± 0.59	0.30 ± 0.08	0.77 ± 0.25	34.74 ± 29.52	1112.6 ± 99.47	20.44 ± 9.10	0.50 ± 0.17	1.81 ± 0.58	5.25 ± 3.36	0.56 ± 0.28	13.70±9.61
CV	7.50	25.54	32.61	85.00	8.94	44.53	33.71	32.00	64.07	49.18	70.13
Max	8.51	0.44	1.19	102.50	1247.0	34.16	0.85	2.68	13.84	1.02	32.50
Min	6.76	0.21	0.36	2.30	961.60	9.16	0.20	1.12	2.29	0.29	5.02
CL (95.0%)	0.45	0.06	0.18	21.90	68.70	6.68	0.13	0.38	2.49	0.20	7.08
Kurtosis	-0.50	-0.87	- 0.70	1.60	-1.26	-1.66	1.27	-1.38	4.10	- 1.41	0.23
Skewness	-0.76	0.26	0.23	1.37	-0.06	0.13	0.43	-0.02	2.00	0.72	1.23
<i>p</i> value (2-tailed)	< 0.0001	< 0.0001	< 0.0001	0.076	< 0.0001	0.001	< 0.0001	< 0.0001	0.000	< 0.0001	0.001
Cluster-IV											
Mean±SD	8.21 ± 0.26	0.37 ± 0.09	0.83 ± 0.51	90.23 ± 43.02	1626.35 ± 324.97	23.33±12.85	0.67 ± 0.25	1.79 ± 0.26	4.94 ± 2.30	0.73 ± 0.20	11.11 ± 7.30
CV	3.20	24.27	61.15	47.67	19.98	55.07	37.28	14.61	46.52	27.46	65.74
Max	8.70	0.51	0.88	164.4	2259.20	45.00	0.95	2.19	9.70	1.13	28.20
Min	7.93	0.28	0.51	38.90	1332.10	7.50	0.19	1.48	2.78	0.46	6.17
CL (95.0%)	0.26	0.09	0.11	39.10	306.40	12.70	0.24	0.26	2.22	0.20	7.27
Kurtosis	0.28	- 0.87	6.88	-0.55	0.88	-0.52	0.59	-1.16	1.93	1.46	5.60
Skewness	0.75	0.80	2.58	0.56	1.32	0.56	-1.06	0.26	1.46	0.86	2.28
<i>p</i> value (2-tailed)	0.001	0.001	0.001	0.022	0.001	0.022	0.001	0.001	0.001	0.001	0.004

*Significant (p < 0.05) **Highly significant (p < 0.01)

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 Table 2
 Pearson correlation coefficient matrix of the 11 elements and their physiochemical characteristics measured for the soils

	рН	EC	OC	P ₂ O ₅	K ₂ O	S	В	Cu	Fe	Zn	Mn
рН	1										
EC	0.707 (**)	1									
OC	0.016	0.070	1								
P_2O_5	0.382 (**)	0.398 (**)	0.053	1							
K ₂ O	0.591 (**)	0.791 (**)	0.178	0.470 (**)	1						
S	-0.101	0.000	0.322 (**)	0.092	0.234 (*)	1					
В	0.098	-0.052	0.033	0.136	0.161	0.191	1				
Cu	-0.023	-0.084	0.440 (**)	0.090	0.168	0.268 (*)	0.124	1			
Fe	-0.744 (**)	-0.561 (**)	0.267 (*)	-0.202	-0.360 (**)	0.191	-0.066	0.419 (**)	1		
Zn	0.052	0.069	0.030	0.441 (**)	0.118	0.065	0.045	-0.076	0.023	1	
Mn	-0.707 (**)	- 0.568 (**)	0.141	-0.328 (**)	-0.398 (**)	0.324 (**)	0.021	0.198	0.639 (**)	-0.059	1

*Significant (p < 0.05)

**Highly significant (p < 0.01)

Table 3 Semivariogram models (ordinary kriging) of 11 soil properties at Kaligaudanahalli microwatershed

Soil properties	Spatial distribution model	Neighbors	Range (m)	Nugget	Partial sill	Sill	Nugget/Sill	RMSP
pН	Exponential, strong	20	2111.35	0.17	0.79	0.95	0.17	1.0
EC (dS m ⁻¹)	Exponential, strong	20	5043.93	0.00	0.014	0.02	0.00	1.2
OC (%)	Exponential, pure nugget	19	683.06	0.00	683.06	683.06	0.00	1.0
P_2O_5 (kg ha ⁻¹)	Exponential, pure nugget	19	750.01	0.00	2763.18	2763.18	0.00	0.9
K ₂ O (kg ha ⁻¹)	Exponential, strong	20	5198.85	60,120.4	183,666.0	243,786.4	0.25	1.0
S (mg kg ⁻¹)	Exponential, pure nugget	19	755.70	0.00	88.97	88.97	0.00	0.9
B mg kg ⁻¹)	Exponential, strong	12	487.66	0.01	0.08	0.10	0.14	1.0
Cu (mg kg ⁻¹)	Exponential, strong	19	1116.00	0.10	0.36	0.46	0.22	1.0
Fe (mg kg ⁻¹)	Exponential, strong	20	2616.47	5.61	25.80	31.41	0.18	1.0
Zn (mg kg ⁻¹)	Exponential, weak	12	503.28	926.78	172.88	1099.66	0.84	0.9
Mn (mg kg ⁻¹)	Exponential, strong	20	2385.61	14.31	163.37	177.69	0.08	1.1

Spatial distribution (S—strong spatial dependence; M—moderate spatial dependence; W—weak spatial dependence; Pure Nugget—no spatial correlation), and spatial distribution model. RMSP—root-mean-square standard prediction error

 $K_2O(0.591^{**})$ and negatively correlated with Fe (-0.744^{**}) and Mn (-0.707^{**}). The results further show that OC has a positive correlation of with S, Cu and Fe, respectively (Table 2).

3.3 Geostatistical analysis

3.3.1 Ordinary kriging

In ordinary kriging, semivariogram models of exponential method were used and presented in Table 3. The data shows that spatial dependence of soil variable is varied from 487.66 m for B to 5198.85 m for available K₂O. In agreement with findings of [39], soil OC, P₂O₅ and S are spatially correlated to distance of 603.06 m, 750.01 m and 755.70 m; however, it showed a small/zero nugget variance. It is observed that the nugget semivariance of OC and P_2O_5 is very small and approached to zero. In the study site, relative nugget effect of pH, EC, K_2O , B, Cu, Fe, Zn and DTPA-extractable Mn is less than 25% showing a strongly spatial dependence. Relative nugget effect for the pH was 17%, Cu-22%, and Fe-18%, indicating a random structured variability and for DTPA-extractable Zn with nugget-to-sill ratio of 84%, indicating weakly structured variability.

3.3.2 Cokriging

Soil pH is affected by land use and management, vegetation type impacts soil pH, a loss of organic matter removal of soil minerals when crops are harvested, erosion of the surface layer, climate, mineral content, but soil texture cannot be changed. In cokriging, the exponential semivariogram models are worked out using pH as ancillary variable and have strong relation with EC, P_2O_5 , K_2O , Fe and Mn. The nugget-to-sill ratio is varied from 0.07 for Mn and 0.16 for P_2O_5 (Table 4) with RMSP of 0.9–1 except for Fe.

3.4 Cross-validation

3.4.1 Ordinary kriging

The cross-validation of results with the calculation of ME, RMSE, AKSE, RMSP, MSPE, R^2 and goodness of fit is presented in Table 5. The lowest root-mean-square error (RMSE) is reported for EC (0.070, 20 neighbors), for B (0.289, 12 neighbors) and for organic carbon (0.318, 19

Table 4	Cross-semivariogram models (cokriging) and parameters for selective soil properties at Kaligaudanahalli microwatershed	
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Soil properties	Spatial distribution model	Range (m)	Nugget	Partial sill	Sill	Nugget/Sill	RMSP
pH versus EC (dS m ⁻¹)	Exponential, strong	2167.64	0.15	1.62	1.77	0.09	0.9
pH versus P ₂ O ₅ (kg ha ⁻¹)	Exponential, strong	1087.68	0.24	1.26	1.50	0.16	0.9
pH versus K ₂ O (kg ha ⁻¹)	Exponential, strong	1490.61	0.23	1.27	1.50	0.15	1.0
pH versus Fe (mg kg ⁻¹)	Exponential, strong	2098.45	0.16	1.60	1.76	0.09	1.1
pH versus Mn (mg kg ⁻¹)	Exponential, strong	1749.57	0.12	1.56	1.67	0.07	0.9

Spatial distribution (S—strong spatial dependence; M—moderate spatial dependence; W—weak spatial dependence; Pure Nugget—no spatial correlation), and spatial distribution model; RMSP—root-mean-square standard prediction error

Soil properties	Neighbors	ME	RMSE	AKSE	RMSP	Regression function	MSPE	R ²	Correlation coefficient
рН	20	0.0030	0.6398	0.66161	0.9746	Y=0.525 * x+3.567	0.0021	0.482	-
EC (dS m ⁻¹)	20	-0.0002	0.0708	0.0597	1.1847	Y=0.511 * x+0.090	-0.0076	0.477	0.707**
OC (%)	19	-0.0001	0.3247	0.3188	1.0196	Y=0.161 * x+0.598	7.1452	0.105	-
P_2O_5 (kg ha ⁻¹)	19	-0.0462	38.854	44.076	0.8910	Y=0.357 * x+20.33	-0.0007	0.293	0.382**
K ₂ O (kg ha ⁻¹)	20	- 3.083	318.71	306.14	1.0395	Y=0.310 * x+432.9	-0.0087	0.380	0.591**
S (mg kg ⁻¹)	19	-0.1358	7.3823	7.8863	0.9363	Y=0.287 * x+13.101	-0.0162	0.297	-
B (mg kg ⁻¹)	12	-0.0015	0.2889	0.2994	0.9699	Y=0.217 * x+0.385	-0.0049	0.138	-
Cu (mg kg ⁻¹)	19	9.9952	0.5431	0.5557	0.9794	Y=0.263 * x+1.273	0.0005	0.238	-
Fe (mg kg ⁻¹)	20	-0.0108	3.5740	3.6237	0.9869	Y=0.540 * x+3.468	-0.0019	0.461	-0.744**
Zn (mg kg ⁻¹)	20	0.1708	31.019	34.476	0.9088	Y = -0.011 * x + 3.755	0.0046	0.004	-
Mn (mg kg ⁻¹)	20	-0.0689	8.7940	7.8318	1.1240	Y=0.592 * x+7.084	-0.0034	0.406	-707**

Table 5 Cross-validation (ordinary kriging) RMSE for all soil properties under study in Kaligaudanahalli microwatershed

ME mean error, *RMSE* root-mean-square error, *AKSE* average kriging standard error, *RMSP* root-mean-square standard prediction error, *MSPE* mean standardized prediction error

Table 6	Cross-validation (cokriging) RMSE fo	or selective soil	properties u	under study in	Kaligaudanahal	li microwatershed
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Soil properties	Neighbors	ME	RMSE	AKSE	RMSP	Regression function	MSPE	R ²	Good- ness of fit
pH versus EC (dS m ⁻¹)	20	-0.008	0.530	0.568	0.934	Y=0.664 * x+2.473	-0.011	0.65	1
pH versus P ₂ O ₅ (kg ha ⁻¹)	19	-0.002	0.599	0.609	0.954	Y=0.579 * x+3.115	-0.010	0.55	1
pH versus K ₂ O (kg ha ⁻¹)	20	-0.006	0.565	0.552	1.054	Y=0.658 * x+2.501	0.023	0.60	1
pH versus Fe	20	-0.007	0.525	0.563	0.891	Y=0.621 * x+2.823	-0.028	0.66	1
pH versus Mn	20	-0.010	0.603	0.533	1.155	Y=0.673 * x+2.397	0.014	0.55	1

ME mean error, *RMSE* root-mean-square error, *AKSE* average kriging standard error, *RMSP* root-mean-square standard prediction error, *MSPE* mean standardized prediction error

neighbors). The root-mean-square standardized prediction error (RMSP) suggests the same since it is less than 1 for pH and DTPA-extractable Zn except EC and DTPAextractable Mn which shows underestimating the variability predictions. The ME and MSPE are negative for EC, OC, K_2O and Mn with AKSE greater than the RMSE, indicating larger kriging variance.

3.4.2 Cokriging

The result revealed that the ME and MSPE are negative with AKSE greater than the RMSE (Table 6). The R^2 value is varied from 0.55 for Mn, P_2O_5 and 0.65 for EC with goodness of fit is 1. Figure 4 shows the scatter plot of estimated versus measured of pH versus EC, P_2O_5 , K_2O and DTPAextractable Fe and Mn data using cokriging methods. Generally, scatter plot is appropriate for quality assurance and accuracy of predictions. It is also useful when there are large numbers of sample points and can provide information about the strength of a relationship between two variables. Based on Fig. 4, all the scatter plots confirm the results of RMSE (Table 6).

4 Discussions

The Kaligaudanahalli microwatershed have crest/summits, back slopes and then foot slopes of granitic terrain supporting medium to deep red soils with five surface textures such as clay (13.1% of total area), sandy clay (23.46%), sandy clay loam (37.54%), sandy loam (13.1%) and loamy sand (4.03%) out of total area [34]. The mean annual rainfall is 670-890 mm, of which about 55-75% is received during kharif season. Descriptive statistical summary (Table 1) shows that these soils are slightly acid to strongly alkaline with low coefficient of variation (< 15%), but have high coefficient of variation for electrical conductivity (> 35%) and for organic carbon, available phosphorus and potassium. This observation is in agreement with the findings of earlier researchers [6, 9, 11]. Considering the critical level of S (10 mg kg⁻¹) and B (0.5 mg kg⁻¹), the high variability of S and B is observed. There is a significant positive relation of organic carbon with available sulfur $(mg kg^{-1}) (R^2) = 0.142^{**}$, significant at 1% level. The available S is the function of organic carbon as reported by Kour and Jalali [20].

It was reported that there is wide spread deficiencies of S, B and Zn in drylands of Karnataka [31, 38, 41], but the soils in the microwatershed are having medium status of S and above critical limit of B status. Similar kinds of observations were reported in various parts of Karnataka [24, 33, 35, 42]. Among all micronutrients, Zn is deficient with high variability. Classical statistics did not show the strongly patchy distribution of some soil parameters and provided mean values that produced medium and large CV for all the soil properties except pH. These findings are in agreement with the results of Cambardella and Karlen [8] and Geypens et al. (1999) [15].

The cluster analysis was done to find out affinity among the 11 soil properties under study and brought out four clusters (Fig. 3). The cluster analysis clearly shows distinct variations in soil properties both within cluster and between the clusters. It is pointed here that the neutral soils grouped under cluster-II showed high variability of organic carbon (CV of 43.99%) with low status of available P₂O₅, below critical limit of DTPA-extractable Zn and Boron (Table 2). It is further observed that soil properties under study have high mean ± SD for cluster-IV as compared to cluster-II. In case of DTPA-extractable Fe and Mn, the trend is opposite to that of cluster-II and IV. Most of the grid points in cluster-II are near to the depressional zones and subjected to seasonal water stagnation, favoring the high concentration of redox sensitive elements such as Fe and Mn. The concentration of Fe and Mn in top soils is negatively related with pH and expressed its exponential relation as:

Eq 10. DTPA-extractable Iron (mg kg⁻¹) = 8.568 e^{-0.02x}, (R^2) = 0.5547**

Eq 11. DTPA-extractable Manganese (mg kg⁻¹) = 8.460 $e^{-0.008x}$, (R^2) = 0.4999**

This equation is significant at 1% level with coefficient of determination (R²) of 0.56** and 0.49** for Mn. The performance of semivariogram are depended dividing of variogram that computed by the function of the distance and characterizes the spatial continuity between points. Five main factors relate to the quality of the data are: distribution, isotropism and anisotropism, variance and range, accuracy, spatial correlation and other factors, and secondary variables [22]. The changing of variogram could be a variable of some parameters such as nugget, partial sill or range, and anisotropy of the data that is also correlated with the distance of the location and the soil properties at a point. In ordinary kriging, exponential semivariograms shows a range of 487 m for B and 5198 m for K_2O (Table 3). Thus, EC and K_2O had a range of more than 5000 m, indicating that both values influenced by neighboring values of soil over greater distances. The number of closest samples chosen varied from 12 for Boron and Zn. The cross-validation of results showed that the lowest root-mean-square error (RMSE) is reported for EC (0.0708, 20 neighbors), for B (0.288, 12 neighbors) and for organic carbon (0.324, 19 neighbors). The root-mean-square standardized prediction error (RMSP) suggests the same since it is less than 1 for pH and DTPA-extractable Zn, but for



Fig. 3 Dendrogram using a hierarchical agglomerative method of complete linkage grouping derived from a Gower's similarity coefficient



Fig. 4 Scatter plot of estimated versus measured and normal QQ-plot of standardized estimation errors with cokriging interpolation of pH versus EC, P₂O₅, K₂O, Fe and Mn



Fig. 5 Thematic maps on available P_2O_5 , K_2O , DTPA-extractable Fe and Mn (cokriging)

underestimating in case of EC and DTPA-extractable Mn as RMSP > 1. The predictions are relatively unbiased for pH, EC, DTPA-extractable Zn and Mn. Since the average kriging standard error (AKSE greater than the RMSE) indicates that the kriging variance is larger than the true estimation variance and indicates that the variogram model is overestimating the variability of the predictions.

Geostatistical-variogram analysis and spatial interpolation (Ordinary kriging and Co-kriging) methods using the threshold limits of each nutrient as per Arora [1] under study. RMSE and QQ-plots (Fig. 4) together with the sum and average errors were considered to be the best estimation methods (Tables 6). The thematic maps show the nutrient zoning of pH versus available phosphorus (P_2O_5), K₂O, DTPA-extractable Fe and Mn under cokriging as best fit model (Fig. 5). The results showed that the area under low status of available phosphorus using cokriging is estimated as 434 ha area (70.86%) with moderately acidic (pH 5.5) to strongly alkaline (pH 9.0) and concentrated in the maximum part of the microwatershed, whereas 178 ha (29.14%) area covered by medium available P_2O_5 with neutral (pH 6.5) to strongly alkaline (pH 9.0) in the microwatershed. Similar exercise was done for deriving thematic maps using cokriging on available K₂O, DTPA-extractable Fe and Mn (Fig. 5). The map shows that 6.75% of moderately acid to neutral soils have medium K₂O, whereas the remaining 93% of moderately acid to strongly alkaline soils have high available K₂O; 26.56% of total area under neutral to strongly alkaline conditions are deficient and 73.44% are sufficient (pH moderately acid to strongly alkaline) in DTPA-extractable Fe. Irrespective of soil pH, the DTPAextractable Mn status is sufficient in these soils. The results are more encouraging for cokriging over ordinary kriging. In practice, identifying the different nutrient zones with respect to soil reaction class is useful for effective soil testbased recommendations and to improve productivity of crops grown in the watershed.

5 Conclusions

The systematic grid survey at 89 locations in Kaligaudanahalli microwatershed was carried out with the aim of deriving maps and spatial variability of soil nutrients for soil-test-based fertilizer recommendations. We employed cluster technique and derived four clusters. The normality test was performed using Kolmogorov–Smirnov test and found that available P₂O₅ is normally distributed with p value of 0.348 and positively correlated with pH ($r=0.382^{**}$). The ordinary and cokriging techniques used to derive thematic maps of available P₂O₅, K₂O, DTPAextractable Fe and Mn. The results showed that the maps derived from cokriging is good for practical recommendations based on soil test values for locally grown crops in the microwatershed. The study further stated that the applications of cokriging technique for deriving nutrient zones are quick and useful for optimizing nutrient plans at watershed level.

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Compliance with ethical standards

Conflict of interest The authors express that there are no conflict of interest with the work contained in this manuscript.

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